

Article

# Identifying the Presence of AMD-Derived Soil CO<sub>2</sub> in Field Investigations Using Isotope Ratios

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Academic Editor: Mostafa Fayek

Received: 8 December 2015; Accepted: 26 February 2016; Published: 3 March 2016

**Abstract:** Recent incidents of hazardous accumulations of CO<sub>2</sub> in homes on or adjacent to reclaimed mine land have been shown to be linked to neutralization reactions between acidic mine drainage and carbonate material. An efficient and economic method is necessary to identify the presence of acid mine drainage- (AMD-) derived CO<sub>2</sub> on reclaimed mine land, prior to construction. One approach to identify the presence of AMD-derived CO<sub>2</sub> is to characterize stable carbon isotope ratios of soil CO<sub>2</sub>. To do so, a viable method is necessary to acquire soil gas samples for isotope ratio analysis. This paper presents preliminary investigations of the effectiveness of two methods of acquiring gas samples (sampling during soil flux measurements and using slam bar) for isotope analysis. The results indicate that direct soil gas sampling is cheaper and provides better results. Neither method is adequate without accounting for temporal effects due to changing gas transport mechanisms. These results have significant implications for safe post-mining land uses and future investigations of leakages from geologic carbon sequestration sites.

**Keywords:** carbon dioxide; isotope ratio; delta-13; soil gas sampling

## 1. Introduction

In the last decade, CO<sub>2</sub> from neutralization reactions between acid mine drainage (AMD) and carbonate minerals has become an emerging geo-hazard. AMD is an acid leachate produced when sulfide minerals are exposed to water and oxygen, resulting in a low pH discharge with elevated dissolved metals and equally high sulphate content. Carbonates, either naturally occurring or added as a mining practice, mitigate the environmental impacts of AMD. Several instances of elevated concentrations of AMD-derived CO<sub>2</sub>, sometimes in excess of 20%, in structures built on or near reclaimed mine land have been reported in the literature [1–4]. High concentrations of CO<sub>2</sub>, usually accompanied by low concentrations of oxygen (O<sub>2</sub>), can lead to asphyxiation. CO<sub>2</sub> concentrations above 3% can cause headaches, sweating, rapid breathing, increased heartbeat, shortness of breath, dizziness, mental depression, visual disturbances, or shaking in humans. Above 10%, CO<sub>2</sub> can lead to unconsciousness and even death. Researchers have reported four fatalities in one instance [2], while five pets were reported killed in another [4].

High concentrations of CO<sub>2</sub> in reclaimed mine soil with AMD and mineral carbonates is not a new phenomenon. For instance, Cravotta *et al.* [5] reported concentrations over 16% 11 m below the surface in reclaimed coal mine spoil. However, until the early 2000s no cases of hazardous concentrations of

CO<sub>2</sub> in homes due to AMD were reported in the literature. In such homes, episodes of sudden rise in CO<sub>2</sub> concentrations in the home atmosphere occur with low pressure atmospheric systems, and, unmitigated, render the home uninhabitable and dangerous [3,4]. Forensic geochemistry, using stable carbon isotopes of CO<sub>2</sub>, has demonstrated that the large quantities of CO<sub>2</sub> that migrate into these homes come from carbonate neutralization reactions in the subsurface and not from biogenic sources like plant root respiration and microbial activity [3].

Natural substances have specific <sup>13</sup>C/<sup>12</sup>C ratios due to fractionation, which is a process of local enrichment or depletion of the heavier isotope (<sup>13</sup>C) depending on their geographical, chemical, or biological origin. This shift in the isotopic ratio of a substance of interest from that of a standard reference material, can be used to identify the origin of a substance. Laughrey and Baldassare [3] successfully used isotope ratios to identify the source of AMD-derived CO<sub>2</sub> in several homes using the δ<sup>13</sup>C-CO<sub>2</sub> (defined by Equation (1); where *R*<sub>sample</sub> and *R*<sub>standard</sub> are the <sup>13</sup>C/<sup>12</sup>C isotope ratios of the CO<sub>2</sub> in the sample and the standard, respectively). In their work (as well as this work), the standard is the Vienna Pee Dee Belemnite reference standard. Due to depletion of <sup>13</sup>C during photosynthesis, biogenic CO<sub>2</sub> is depleted in <sup>13</sup>C or enriched in <sup>12</sup>C with δ<sup>13</sup>C (‰) in the range of −35‰ to −8‰. Carbonate rocks are relatively enriched in <sup>13</sup>C due to isotope equilibrium exchange reactions in the inorganic carbon system (atmospheric CO<sub>2</sub>-dissolved bicarbonate-solid carbonate) with δ<sup>13</sup>C (‰) in the range of −7‰ to +4‰ [3,6,7].

$$\delta^{13}\text{C}_{\text{sample}} = \left( \frac{R_{\text{sample}} - R_{\text{standard}}}{R_{\text{standard}}} \right) 1000 \quad (1)$$

It is important to note that the studies, prior to this work, concentrated on the gases intruding into structures (mainly homes) built on or adjacent to reclaimed mine land. Due to accumulation in the basements of these structures, indoor air sampling was relatively easy. In order to investigate this phenomenon, prior to residential or commercial real estate construction, there is a need for soil gas samples (or some sort of sampling indicative of subsurface CO<sub>2</sub>). Soil gas sampling for isotopic studies face two challenges: (i) differences in diffusion rates of the isotopologues, <sup>13</sup>CO<sub>2</sub> and <sup>12</sup>CO<sub>2</sub>, through soil [8]; and (ii) soil transport mechanisms that drive CO<sub>2</sub> flow to the surface [8]. It has been shown that CO<sub>2</sub> from AMD-carbonate reactions accumulate at depth and only flow to the surface in episodes [4,5]. In addition, any sampling method needs to be fast and economic to cover the large tracts of land under study. Hence there is a need to evaluate different methods of acquiring soil gas samples for isotope ratio mass spectrometry, with a view to identify potentially hazardous CO<sub>2</sub>, delineate the extent of such hazards, and apportion mitigation liability.

The two sampling methods investigated in this work can be classified as whole air-active and air-passive methods (using language in ASTM D5314 [9]). In the whole air-active method, the researchers used a slam bar to drive a small diameter hole (~9.5 mm) into the soil, creating a low pressure zone, and draw a bulk sample of soil gas using a hand aspiration pump. In the whole air-passive method, samples of soil atmosphere migrating into an accumulation chamber (during accumulation chamber flux measurement) are drawn using a syringe. The air-passive method done in conjunction with accumulation chamber flux measurement is relevant because the authors were also evaluating the potential to use soil flux monitoring to delineate the hazard posed by AMD-derived CO<sub>2</sub> [10].

The objective of this work was to evaluate the effectiveness of the described sampling methods for identifying the presence of hazardous AMD-derived CO<sub>2</sub> on reclaimed mine land, prior to commercial or residential real-estate development. The research team conducted two sampling campaigns on two different reclaimed mine plots in the United States (US). The sampling methods were evaluated based on speed of sample acquisition, number of soil atmosphere samples required to obtain a single estimate of soil δ<sup>13</sup>C-CO<sub>2</sub>, and the ability to predict underlying AMD-derived CO<sub>2</sub>. On both sites, δ<sup>13</sup>C-CO<sub>2</sub> from monitoring wells was used to determine the ability of the sampling method to predict underlying AMD-derived CO<sub>2</sub>. Both plots of land were known to have AMD-derived CO<sub>2</sub> in the sub-soil [3,4].

Safe post-mining land use requires an ability to evaluate hazards. The emergence of AMD-derived CO<sub>2</sub> as a geohazard requires effective and economic methods to evaluate such hazards. There is also a need to identify the source of any anomalously high CO<sub>2</sub> concentrations to establish liability for mitigation efforts, if necessary. This work will provide valuable insight into the potential pitfalls and benefits of the evaluated sampling methods. The results will encourage further research to develop more robust sampling methods for identifying AMD-derived CO<sub>2</sub> on reclaimed mine land.

## 2. Soil Gas Sampling Methods

Soil gas sampling methods are used in various applications including agriculture, volcanology, and environmental site characterization to monitor various trace gases and contaminants [11–13]. Generally, soil gas sampling methods used for monitoring the vadose zone can be classified into active and passive methods, based on how a gas sample is obtained, or whole-air and sorbent methods, depending on whether all or only a portion of the constituents of the soil atmosphere is collected [9]. Active methods are those methods that use a device in the subsurface to withdraw a soil gas sample while passive methods allow the atmosphere within the sampling device to come into equilibrium with the soil atmosphere. Whole-air methods withdraw a sample of the entire soil atmosphere for analysis, while sorbent methods sample only contaminants adsorbed onto a collection medium. The particular soil gas sampling method suitable for any application depends on several factors including the type of sampling system, the methodology for applying the system and the quality assurance/quality control (QA/QC) protocol [9]. Other considerations include cost, speed and ease of sample acquisition. ASTM D5314-92 [9] provides standard protocols for monitoring soil gases in the vadose zone.

In whole air-active sampling, a bulk sample of the soil atmosphere is forced into a collection device through a probe or similar device. The sample of soil gas is then sent to the laboratory for analysis or coupled with a field analysis apparatus for analysis. In some instances, the collection device (or probe) is driven into the soil with a drill or similar device. In other instances, the collection device is such that researchers have to use a slam bar to first drive a hole before sample collection [14]. A disadvantage of this method is that the process of driving the hole and its removal, prior to sample collection, can result in venting of soil gases [9]. Therefore, the subsequent sample may not be representative. Also, since the hole and the tubes used in sampling have to be purged before gas collection, the soil atmosphere is further purged and can contribute to non-representative samples. Nevertheless, soil sampling using a slam bar and a collection method (such as a sample bag with a pump) is common in the literature [15–17].

In most applications of this method, the goal is to identify and estimate the concentration of contaminant gases in the soil atmosphere [16,17]. However, the same sample can be used in isotope ratio mass spectrometry to determine isotopic signatures. So long as the gas sample is a representative one, no further adjustments are necessary for isotopic analysis.

In contrast to whole air-active sampling, researchers conducting whole air-passive sampling collect a bulk sample of soil atmosphere from a near-surface sampling horizon into a collection device via a flux chamber or other similar apparatus. During accumulation chamber flux measurement, a collar is installed to constrain a section of soil and an accumulation chamber is installed over the soil. Soil gas gradually migrates into the headspace in the chamber. The flux is estimated from measurements of the concentrations of a particular gas constituent over time. A gas sample drawn from the chamber headspace is a whole air-active sample. To assess the isotopic composition of a particular gas constituent, the sample drawn from the chamber, during flux measurement, can be used for isotope ratio mass spectrometry (IRMS).

The isotopic composition of soil gas, determined using a gas sample acquired during chamber accumulation flux measurement, can be affected by gas mixing (between atmospheric and soil gases in the chamber) and diffusion [18].

Gas mixing is an issue because the sample is not a sample of the soil pore gas but rather a mixture of the soil and atmospheric gases. Keeling [19,20] proposed a linear mixing model (Equation (2)),

which can be used to predict the  $\delta^{13}\text{C}$  of  $\text{CO}_2$  in the presence of multiple sources.  $\delta^{13}\text{C}_M$ ,  $\delta^{13}\text{C}_B$ , and  $\delta^{13}\text{C}_S$  are the delta-13 of the  $\text{CO}_2$  mixture, background  $\text{CO}_2$ , and the source  $\text{CO}_2$ , respectively; and  $c_B$  and  $c_M$  are the background and mixture  $\text{CO}_2$  concentrations, respectively. Using Equation (2), one can obtain the  $\delta^{13}\text{C}$  of the source gas graphically as the intercept on the vertical axis by plotting  $\delta^{13}\text{C}$  of the mixture  $\text{CO}_2$  against the inverse of the mixture  $\text{CO}_2$  concentration. This is the so called Keeling plot.

$$\delta^{13}\text{C}_M = c_B \left( \delta^{13}\text{C}_B - \delta^{13}\text{C}_S \right) \frac{1}{c_M} + \delta^{13}\text{C}_S \quad (2)$$

However, the Keeling plot is most useful when the  $\delta^{13}\text{C}$  measurements are independent (e.g., measurements from different sample locations). When all the samples are from the chamber headspace during the same closure, the  $\delta^{13}\text{C}$  measurements are not independent. In such a case, it is possible to estimate the source isotopes directly from the isotope ratios and concentrations of two samples [18]. In this work, we used Equation (3) to estimate the  $\delta^{13}\text{C}$  of the source  $\text{CO}_2$ ,  $\delta^{13}\text{C}_S$ , from the concentration,  $c_M^i$ , and  $\delta^{13}\text{C}\text{-CO}_2$ ,  $\delta^{13}\text{C}_M^i$ , of two samples ( $\{i : i = 1, 2\}$ ).

$$\delta^{13}\text{C}_S = \frac{\left( \delta^{13}\text{C}_M^1 \right) c_M^1 - \left( \delta^{13}\text{C}_M^2 \right) c_M^2}{c_M^1 - c_M^2} \quad (3)$$

Estimating isotopologue composition of  $\text{CO}_2$  from a sample that is drawn after diffusing through soil can be problematic because of the aforementioned difference in diffusion rates of  $^{13}\text{CO}_2$  and  $^{12}\text{CO}_2$  through soil. Fortunately, it has been shown theoretically and experimentally that, in many instances, the different diffusion rates result in a 4.4‰ difference in the  $\delta^{13}\text{C}$  [7]. This knowledge can be used to adjust estimates of isotopic ratios determined from samples which diffuse, preferentially, through soil (i.e., not in equilibrium).

### 3. Materials and Methods

#### 3.1. Study Sites

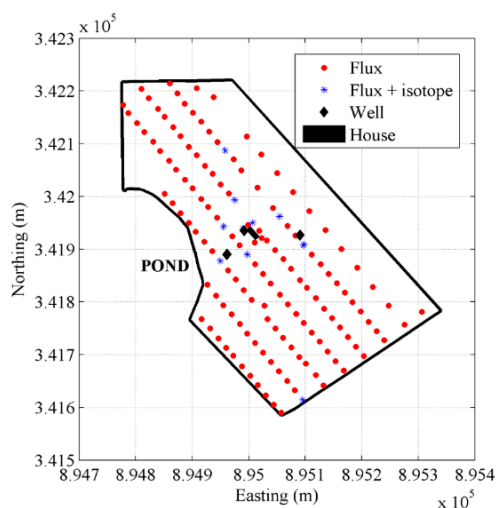
##### 3.1.1. The Hudson Site

This site is located in Pike County in south-western Indiana ( $38^\circ 19' 42''$  E and  $87^\circ 08' 27''$  N). It is a reclaimed surface coal mine that has a single story residential building with a walk-in basement constructed on it. The home has been the subject of scientific and regulatory investigations of stray  $\text{CO}_2$  gases since 2006 due to intermittent episodes of elevated concentrations of stray  $\text{CO}_2$ . Five pets have died in the home and one teenager has been sent to the emergency room as a result of episodes of high  $\text{CO}_2$  concentrations in the basement and garage of the home [4].

Mining was carried out between 1986 and 1992 and the site was reclaimed with lime amendment and about 0.91 m of top soil capping. The spoil material extends about 11.6 meters below the surface. The area receives an average total rainfall of 1184.1 and 304.8 mm of snowfall, annually. The average daily temperature is  $12.7^\circ\text{C}$  with minimum and maximum temperatures of  $-6.1$  and  $31.1^\circ\text{C}$  for the winter (January) and summer (July) months, respectively [21].

The predominant plant cover is  $\text{C}_3$  type (native grasses). We did not evaluate the isotopic signature of organic matter because we assumed this will not have a significant bearing on the results. However, such analysis may be warranted in cases where there is significant  $\text{C}_4$  plant cover.

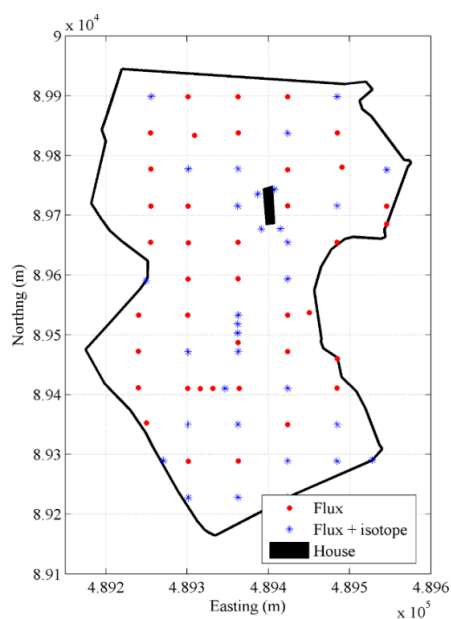
The area under investigation (Figure 1) covers about 16 hectares with soils described as Fairpoint loam. The site was reclaimed at  $1^\circ$  to  $15^\circ$  slopes [22].



**Figure 1.** Hudson site showing flux and isotope sampling locations. All coordinates are in Indiana State plane coordinates.

### 3.1.2. The Godin Site

Figure 2 shows the location of the Godin site. This site also has a single family residential building, which has had episodes of elevated  $\text{CO}_2$  concentrations due to AMD-derived  $\text{CO}_2$  in the subsoil. The home was built on the spoil of the reclaimed Godin Mine, which is about 21.34 m thick. The site is situated near Jenners, a town in Somerset County, southwestern Pennsylvania ( $40^\circ 08' 2''$  E and  $79^\circ 02' 52''$  N). Stray  $\text{CO}_2$  in the Godin residence was investigated by the Pennsylvania Department of Environmental Protection (PA-DEP) in 2003 [3]. The mine's permit required an operational plan that included disposing of the material from the surface mine pit floor in pods at least 3.05 m above the pit floor. Lime was added to the pit floor at a rate of 45 tonnes/hectare prior to backfilling and grading. The total annual precipitation averages 1053 and 881 mm of rainfall and snowfall, respectively. The average daily temperature is  $6.7^\circ\text{C}$  and ranges from an average minimum temperature of  $-4.9^\circ\text{C}$  to a maximum of  $18.9^\circ\text{C}$  for the winter (January) and summer (July) months, respectively [21].



**Figure 2.** Godin site showing flux and isotope sampling locations. All coordinates are in Pennsylvania State plane coordinates.

Similar to the Hudson Site, the vegetation covering this site is also predominantly grasses ( $C_3$  plants). We did not evaluate the isotopic signature of organic matter for the same reasons discussed above.

### 3.2. Sampling Methods

This section describes the two sampling methods used in this research to acquire soil gas samples for laboratory and subsequent analysis. All gas samples were subjected to the same gas chromatography (GC) and IRMS techniques to estimate the  $\delta^{13}C$  of the  $CO_2$  in the sample. The researchers used Isotech Laboratories, Inc. (Champaign, IL, USA, [www.isotechlabs.com](http://www.isotechlabs.com)), a reputable commercial laboratory with experience doing forensic geochemistry for the oil and gas and mining industries, for all the laboratory analysis. The two methods used to acquire soil gas samples are: (1) an accumulation chamber technique, which is a whole air-passive method; and (2) a slam bar and hand aspiration technique, which is a whole air-active method. These methods are described in the subsections below.

Admittedly, it would have been better (from a scientific experimental design standpoint) to have used both methods on both sites to ensure comparable results. However, the practical limitations of the research project (The initial research design included only the accumulation chamber technique which the team used at the Hudson site. Upon reviewing the results, the team decided to use the slam bar method at the Godin site as an alternative. At this stage, the residents at the Hudson site had dropped out of the study and we did not have the funding to conduct further isotope tests.) did not allow this. Instead, the research team used the accumulation chamber at the Hudson site and the slam bar and hand aspiration technique at the Godin site. This is a limitation of the current research that will need to be addressed with future research. However, the authors believe the work presented here makes enough of a contribution to the literature on its own.

#### 3.2.1. Accumulation Chamber Technique

This technique involved collecting gas samples from the headspace of an accumulation chamber for isotope analysis during flux sampling. In this work, researchers used a portable automated accumulation chamber monitor (model: LI-8100) manufactured by Licor Biosciences Inc. (Lincoln, NE, USA). The chamber diameter is 20 cm with a headspace volume of 4.8 L. The LI-8100 monitor has an infra-red gas analyzer (IRGA), which is used to measure the  $CO_2$  concentration during flux measurement. The monitor has tubes that draw air from the chamber headspace to the IRGA. In this application, a T-fitting with a septum is installed in the tubes to facilitate sample acquisition. Researchers used a syringe to draw a sample from the air stream through the septum.

It is necessary to acquire multiple samples to account for the gas mixing issues discussed earlier. In this work, the authors decided to draw three samples to minimize cost but allow for multiple estimates (each estimate requires two samples). The chamber was deployed for 20 min (as opposed to the usual 2 min for flux measurement) to allow for all three samples to be collected during flux measurement. The first sample was taken after the chamber  $CO_2$  concentration reached at least 500 ppm (according to the LI-8100 monitor), to ensure there is enough  $CO_2$  for isotopic analysis. Depending on the flux at the particular sampling point, this took approximately 10 min. Following this first sample, two subsequent samples were drawn in the remaining time. Each sample was 60 mL (Note that the sample volume need not be 60 mL. Samples of 10 mL or more could suffice and smaller sample volumes are desirable. We used 60 mL samples to ensure duplicate testing was possible without more field visits. We assume that drawing three 60 mL samples over 20 min did not significantly disturb the chamber headspace. We confirmed this by monitoring the  $CO_2$  concentrations during sampling.) and was transferred into the sample bag via a one-way valve with a septum.

Twenty-seven samples were collected from nine sample locations around the property on 2 April, 2010 (Figure 1). All the samples were shipped to Isotech for GC/IRMS analysis. The results of the



isotopic analysis were used to estimate the carbon isotopic ratios of the soil CO<sub>2</sub>, accounting for both preferential diffusion and gas mixing. The results are reported in the Results and Discussions Section.

### 3.2.2. Slam Bar and Hand Aspiration Technique

The second technique involved drilling a hole, about 9.5 mm in diameter, into the soil to a depth of 0.61 m (2–3 ft) with a slam bar. The researcher then inserts a clear, polyethylene tubing of the hand aspiration pump into the soil to about 5 cm short of the bottom of the hole. The top of the hole is sealed with a cap and soil. After purging the system (tubing and pump), a soil gas sample was hand aspirated into a pre-evacuated specialty gas sample bag. A single sample was collected from each of the 32 sample locations at the Godin site on 15 July 2010 (Figure 2).

The samples were then shipped to Isotech (Isotech Laboratories, Inc.) for GC/IRMS analysis. The results are reported in the Results and Discussions Section.

## 4. Results and Discussions

### 4.1. Accumulation Chamber Results

Table 1 shows the results of the laboratory analysis of the 27 samples. The table shows the concentration of CO<sub>2</sub> in the samples and the  $\delta^{13}\text{C}$  of the CO<sub>2</sub>. From the results, we observe that as the concentration of CO<sub>2</sub> increases (more soil gas migrates into the chamber headspace),  $\delta^{13}\text{C}$ -CO<sub>2</sub> becomes more negative. This indicates the ratio of <sup>13</sup>C to <sup>12</sup>C in the soil CO<sub>2</sub> is lower than that in the atmospheric CO<sub>2</sub>. Using Equation (3), we estimated the carbon isotope ratio of the soil CO<sub>2</sub> for each sample location. The results are shown in Table 2. We estimated the  $\delta^{13}\text{C}$  three times using all combinations of the three samples. We consider the estimate using the first and third sample (column 4 of Table 2) the most reliable since it represents the largest difference in CO<sub>2</sub> concentration. Based on these estimates, the average  $\delta^{13}\text{C}$ -CO<sub>2</sub> of the nine sample locations was found to be  $-24.5\%$ .

**Table 1.** Isotope ratio results for Hudson site.

Sample ID	[CO <sub>2</sub> ] (ppm)	$\delta^{13}\text{C}$ -CO <sub>2</sub> (‰)	Sample ID	[CO <sub>2</sub> ] (ppm)	$\delta^{13}\text{C}$ -CO <sub>2</sub> (‰)	Sample ID	[CO <sub>2</sub> ] (ppm)	$\delta^{13}\text{C}$ -CO <sub>2</sub> (‰)
B11-1	621.53	−15.2	C15-1	1090.55	−19.7	D16-1	658.19	−16.7
B11-2	899.10	−18.4	C15-2	1211.28	−20.4	D16-2	699.67	−17.6
B11-3	1029.50	−18.9	C15-3	1321.18	−20.9	D16-3	709.37	−17.4
B14-1	622.32	−15.7	C17-1	601.74	−12.6	E14-1	787.78	−17.4
B14-2	826.19	−18.4	C17-2	715.39	−13.8	E14-2	1049.69	−20.0
B14-3	981.77	−20.2	C17-3	751.47	−14.2	E14-3	1362.63	−21.6
B21-1	638.50	−15.1	D13-1	791.20	−17.6	H3-1	556.25	−12.9
B21-2	670.01	−16.0	D13-2	1036.50	−19.5	H3-2	627.16	−14.0
B21-3	597.56	−14.7	D13-3	1355.50	−21.1	H3-3	654.07	−14.5

**Table 2.** Isotope ratio results using Keeling plot and adjusted for preferential diffusion.

Sample ID	$\delta^{13}\text{C}$ -CO <sub>2</sub> (‰) Samples 1–2	$\delta^{13}\text{C}$ -CO <sub>2</sub> (‰) Samples 2–3	$\delta^{13}\text{C}$ -CO <sub>2</sub> (‰) Samples 1–3
B11	−25.57	−22.35	−24.54
B14	−26.64	−29.76	−27.99
B21	−34.24	−26.72	−20.94
C15	−26.72	−26.41	−26.57
C17	−20.15	−22.13	−20.63
D13	−25.63	−26.30	−26.01
D16	−31.88	−2.97	−26.40
E14	−27.82	−26.97	−23.56
H3	−22.63	−26.15	−23.60
Average	−26.81	−23.31	−24.47
Standard Deviation	4.28	7.98	2.54
Skewness	−0.35	2.52	0.38

#### 4.2. Slam Bar Results

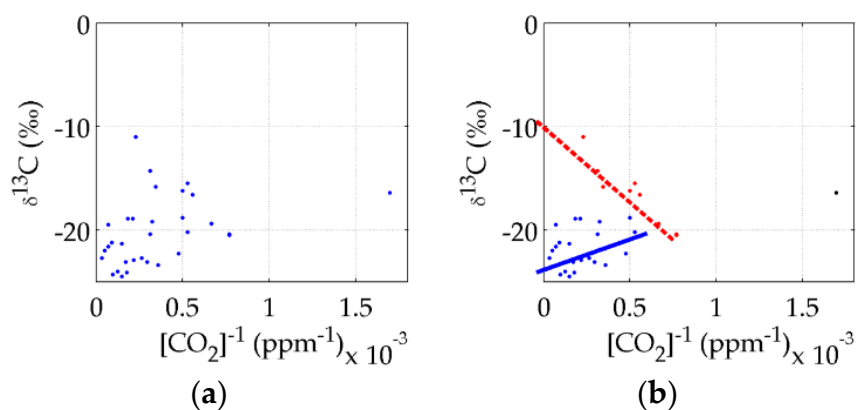
Table 3 shows the results of laboratory analysis of the soil gas samples acquired using the slam bar and hand aspiration pump. The table shows the CO<sub>2</sub> concentrations and δ<sup>13</sup>C-CO<sub>2</sub> of the soil gas samples. The CO<sub>2</sub> concentrations are much higher than day time concentrations in the atmosphere, which averages around 400 ppm [23]. The CO<sub>2</sub> concentrations range from 590 to 29,900 ppm with a mean of nearly 6000 ppm. The high CO<sub>2</sub> concentrations indicate that the samples are not samples of atmospheric gases circulated through the soil.

**Table 3.** Isotope results for Godin site.

Sample ID	[CO <sub>2</sub> ] (ppm)	δ <sup>13</sup> C-CO <sub>2</sub> (‰)	Sample ID	[CO <sub>2</sub> ] (ppm)	δ <sup>13</sup> C-CO <sub>2</sub> (‰)	Sample ID	[CO <sub>2</sub> ] (ppm)	δ <sup>13</sup> C-CO <sub>2</sub> (‰)
A-1	2800	−23.4	C-7	2100	−22.3	E-11	6800	−21.3
A-11	1900	−20.2	C7-100	4400	−11.0	E-4	4600	−22.9
A-6	590	−16.4	C7-50	5700	−24.1	F-11	10,400	−24.3
B-10	14,200	−21.6	C-8	4700	−18.9	F-3	11,200	−21.2
B-12	5800	−23.1	D-11	2900	−15.8	H-1	2000	−18.8
B-3	29,900	−22.7	D-12	14,000	−19.5	H-2	3200	−14.3
B-8	1300	−20.5	D-2	8000	−24.0	H-3	5500	−18.9
BC9-150	2000	−16.2	D-5	1800	−16.6	H-4	19,500	−22.0
C-10	1900	−15.5	D-6	3200	−20.4	Mean	5947	−20.1
C-12	3100	−19.2	D-9	1300	−20.4	Standard Deviation	6197	3.29
C-3	3800	−22.7	E-1	1500	−19.4	Skewness	2.396	0.864
C-4	3400	−23.1	E-10	6800	−24.5	-	-	-

The δ<sup>13</sup>C-CO<sub>2</sub> range from −24.5‰ to −11‰ with a mean of −20.1‰. We did not correct our δ<sup>13</sup>C-CO<sub>2</sub> results for any mixture or diffusion effects. However, it is important for the reader to note that there may be some mixing between CO<sub>2</sub> from the near surface soil horizon and the deeper AMD-related CO<sub>2</sub> during sampling. These two sources do not have the same ratios of CO<sub>2</sub> isotopologues due to differential diffusion of the two isotopologues from the AMD-related CO<sub>2</sub> reservoir to the surface. It is likely, however, that the difference between the ratio of the two CO<sub>2</sub> isotopologues in the AMD-related CO<sub>2</sub> and the sampled CO<sub>2</sub> is lower than 4.4‰ since the tubing is flushed prior to sampling. These factors make it difficult to adjust for the effect of preferential diffusion without further research. Consequently, we chose not to adjust our isotope ratio results.

Given that the slam bar samples are independent, we proceeded to plot Keeling plots of the field (Figure 3). We used the plot to evaluate, in a preliminary fashion, whether any conclusions can be made with regards to the sources of CO<sub>2</sub> at the study site.



**Figure 3.** Keeling plot of isotope results from Godin Site. (a) Raw; (b) With interpretation.



### 4.3. Discussions

To evaluate the usefulness of the two methods for field investigations, the authors focused on three criteria: speed of sample acquisition, number of gas samples required for analysis, and ability to indicate the presence of AMD-derived CO<sub>2</sub>.

The speed of acquiring samples and the number of samples are directly proportional to the cost of the survey. This is because personnel costs increase with time spent acquiring samples and laboratory costs (time, man-power, and consumables) increase with the number of samples.

The experience on this project indicates that the time required to acquire a sample is significantly less for the slam bar method compared to the accumulation chamber method. The actual sample acquisition (taking samples by syringe or hand aspiration pump) is similar; although, the hand aspiration method is slightly faster (no transfer is required as the sample is pumped directly into the sample bag). The major difference between the two methods is the long time required for the soil gas to migrate into the chamber (20 min in this study). Based on the authors' experience, this is possibly the least time required to acquire three samples (although, we note that the estimates can be obtained with only two samples). It took about 10 min for the CO<sub>2</sub> concentration to reach 500 ppm, at which point the first sample can be drawn. The remaining 10 min is just about enough to draw the remaining two samples. In instances where low fluxes resulted in more than 10 min to reach 500 ppm, the research team found itself struggling to acquire all three samples in time. On the contrary, the slam bar method did not require any such waiting period. Researchers were able to draw the sample soon after driving the hole into the soil. The time it takes to drive the hole is not significantly longer than the time it takes to drive the collar and set up equipment for accumulation chamber measurement.

The slam bar method requires only one sample per sampling location. To account for the gas mixing, the accumulation chamber method requires at least two from the same sampling location. In this work, the authors used three samples. The advantage of using only two samples, spaced as far apart as possible to ensure maximum change in CO<sub>2</sub> concentration, is that there is limited disturbance of the chamber headspace atmosphere. However, three samples ensure that researchers have a higher chance of obtaining an estimate if laboratory results for one sample yield an unreliable result. For example, the second sample at D16 appears to be an anomaly. Without the third sample, we would not have been able to obtain a reliable estimate of the delta-13. Regardless of how many samples are taken per chamber closure (two or three), the accumulation chamber method will still require more samples than the slam bar method.

Thus, with the same number of soil gas samples, the slam bar method would provide more  $\delta^{13}\text{C-CO}_2$  estimates. For instance, in this work, the two surveys obtained similar numbers of samples (27 and 32 samples for the accumulation chamber and slam bar methods, respectively). Yet, the slam bar method resulted in three times more  $\delta^{13}\text{C-CO}_2$  estimates (9 and 32, respectively). The higher number of samples allow for statistically more significant estimates of the population mean and nature of the distribution. However, the accumulation chamber approach is useful in situations where isotopic characterization is used in conjunction with flux measurement to delineate the hazard [10].

The ability to accurately indicate the presence of AMD-derived CO<sub>2</sub> is, perhaps, the most important criteria. Table 4 shows data from the authors' other work on the  $\delta^{13}\text{C-CO}_2$  estimates of the underlying CO<sub>2</sub>. As the table shows, the  $\delta^{13}\text{C}$  values of the underlying CO<sub>2</sub> are relatively closer to zero (the standard) than the values observed using both methods (Tables 2 and 3). The mean  $\delta^{13}\text{C-CO}_2$  estimate for the Godin site using the slam bar method is  $-20.1\text{‰}$  (with standard deviation of  $3.29\text{‰}$ ). This compares to  $\delta^{13}\text{C-CO}_2$  estimate of  $-4.2\text{‰}$  for soil gases in the reclaimed mine subsoil (Table 4). The mean for the Hudson site using the accumulation chamber method is  $-24.47\text{‰}$  (with standard deviation of  $2.5\text{‰}$ ) compared to  $-11.4\text{‰}$  to  $-5.0\text{‰}$  for soil gases in the reclaimed mine subsoil (Table 4).

**Table 4.** Isotope ratios of the underlying CO<sub>2</sub>.

Site	Description	Depth (m)	[CO <sub>2</sub> ] (%)	δ <sup>13</sup> C-CO <sub>2</sub> (‰)	Source
Godin	Basement of residence	-	9.51	-4.1	[3]
	Monitoring well	21.3	8.92	-4.2	
Hudson	Monitoring well MW1S	5.79	8.66	-5.0	[24]
	Monitoring well MW2S	5.79	13.30	-5.3	
	Monitoring well MW3S	5.79	9.40	-5.3	
	Monitoring well MW1D	11.58	17.58	-11.4	
	Monitoring well MW2D	11.58	11.24	-9.5	
	Monitoring well MW3D	11.58	0.52	-10.2	

Figure 3 provides further insight on the sources of the CO<sub>2</sub> at the Godin site. At first glance (Figure 3a), there appears to be no discernible trend to the δ<sup>13</sup>C-CO<sub>2</sub> values. Upon careful examination, however, there appears to be multiple sources underlying the data. The two clusters that appear to be the most prominent are identified in Figure 3b. These two clusters seem to point to sources with δ<sup>13</sup>C-CO<sub>2</sub> of -9‰ and -23‰. Although, none of these sources have the same signature as the limestone source, it is conceivable that the -9‰ source may have some contribution from a limestone-based source. This is, however, not conclusive evidence of the presence AMD-derived CO<sub>2</sub> at this study site. It only points to the presence of multiple sources, one of which is likely to be organic matter (-23‰).

We observe then that both methods fail to predict the presence of AMD-derived CO<sub>2</sub> in the subsoil. The authors hypothesize that this is due to the episodic nature of the events resulting from varying transport mechanisms. It is important to note here that the CO<sub>2</sub> in the homes during episodes show δ<sup>13</sup>C-CO<sub>2</sub> values that are similar to those observed in the monitoring wells [3]. Thus, when AMD-derived CO<sub>2</sub> is transported from the subsoil into homes during episodes (of low barometric pressure events) in significant quantities, the isotope ratios are similar. The authors hypothesize that under normal conditions when the CO<sub>2</sub> transport through the soil is controlled by concentration gradients, the amount of AMD-derived CO<sub>2</sub> transported through the soil is not high enough to significantly alter the carbon isotope composition of native biogenic CO<sub>2</sub>. However, during episodes, when the CO<sub>2</sub> transport is controlled by pressure gradients, the transport of significant amounts of AMD-derived CO<sub>2</sub> should alter the carbon isotope composition of CO<sub>2</sub> measured with either method [3,4,8,10]. This hypothesis can be tested by acquiring soil gas samples during episodes, which is beyond the scope of this work but will be pursued as future research.

However, given the advantages of the slam bar method over the accumulation chamber method (higher speed of sample acquisition and lower number of samples required), it appears the slam bar method shows more promise than the accumulation chamber method. Both methods, without accounting for gas transport mechanisms, will however yield false negatives. It is, therefore, important not to rule out this hazard without careful long term monitoring that accounts for pressure pumping effects.

These results and observations have a significant effect on regulatory and legal liability investigations for post-mining land-use of reclaimed mine land (in particular, coal mine land). The results show that using these methods in field investigations to declare a piece of land free of this hazard can be problematic. While positive detection of high fluxes and carbon isotope signatures can be conclusive, the lack of such evidence may not indicate the absence of risks. It is important to consider results such as this in assigning legal liability or declaring a plot of land safe for residential or commercial real-estate use. These observations are also relevant for carbon sequestration sites, as some of these same transport mechanisms may affect CO<sub>2</sub> leakages trapped in subsoil.

## 5. Conclusions and Recommendations

This work investigated the effectiveness of two methods of soil gas sampling (sampling during soil flux measurements and sampling using a slam bar and hand aspiration pump) in acquiring representative samples for stable carbon isotope analysis. The goal of such analysis is to determine, using isotope ratio mass spectrometry, the carbon isotope signature of the soil CO<sub>2</sub> for source identification. The methods were evaluated based on the speed of sample acquisition, number of samples required for the analysis, and ability to accurately identify the presence of acid mine drainage (AMD-) derived CO<sub>2</sub>.

The results indicate that the slam bar method is faster and requires fewer samples per sample location. The accumulation chamber method requires at least two times more samples to cover the same number of sampling locations. This means the slam bar method will be cheaper per sampling location. However, the accumulation chamber method will be useful in situations where accumulation chamber flux measurement is part of the sampling protocol already. Neither method adequately predicts the presence of AMD-derived CO<sub>2</sub> in the subsoil. This is most likely due to the absence of significant transport of AMD-derived CO<sub>2</sub> from the subsoil upwards. The authors hypothesize that this will change if the samples were acquired during episodes of barometric pressure pumping where significant amounts of AMD-derived CO<sub>2</sub> is transported through the soil to the atmosphere. Even so, the slam bar method appears to predict carbon isotope ratios closer to the isotope ratios of the AMD-derived CO<sub>2</sub> than the accumulation chamber method.

These results have significant implications for declaring commercial and residential real-estate as safe post-mining land uses of reclaimed mine land with potential for AMD and future investigations of leakages from geologic carbon sequestration sites. The observations from this work imply false negatives are likely and stakeholders have to take this into consideration when acquiring soil gas samples using either method.

Further work is required to test the hypothesis that these sampling methods will perform better if CO<sub>2</sub> migration episodes due to low barometric pressure are accounted for in the monitoring. This requires long term monitoring of CO<sub>2</sub> fluxes, barometric pressure and other weather-related variables, and isotopic signatures. In addition, future work should examine the effect, if any, of fractionation bias due to differences in CO<sub>2</sub> isotopologue diffusion, especially when using the slam bar method. The effect of any such fractionation should be examined under normal conditions and during CO<sub>2</sub> migration episodes since the effect might differ under the two conditions.

**Acknowledgments:** The authors would like to acknowledge: Office of Surface Mining, Reclamation and Enforcement's Applied Science program's project #S09AC15437 and the Government of Botswana for financial support; Bret A. Robinson for his assistance in accessing the study site; and Bismark Osei for his assistance in this work. The authors are also grateful for the reviewers' comments on the initial manuscript which improved the analysis and discussions in this paper.

**Author Contributions:** Kwame Awuah-Offei conceived the research project, wrote the proposal to obtain funding and directed the researcher effort. He also wrote the initial draft of the manuscript for the co-authors review. Moagabo Mathiba revised the initial experimental design as needed, conducted data analysis, wrote key portions of the manuscript, and reviewed the entire manuscript. Fred Baldassare provided isotope geochemistry expertise, participated in experimental design and proposal development, and reviewed the manuscript. All authors participated in the field data collection.

**Conflicts of Interest:** The authors declare no conflict of interest.

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